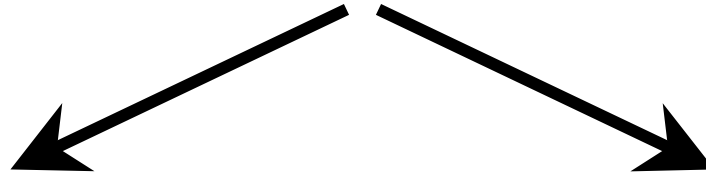


Net primary production (NPP) = f [biomass, physiology, $g(I_0)$]



$NPP = f$ [chlorophyll, Z_{eu} , P_{opt}^b , $g(I_0)$]

{0.02 to 30 mg m⁻³}

{<10 to >120 m}

{0 to 1}

{2 to 300 mg m⁻²}

{0 to 30 mg C mgChl⁻¹ h⁻¹}

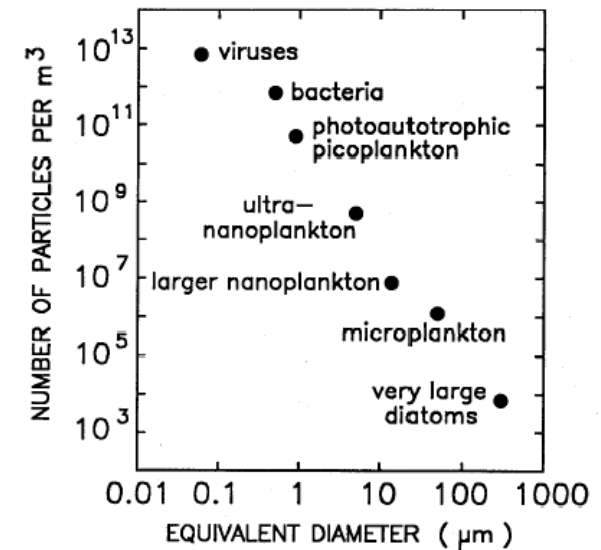
$NPP = f$ [Carbon, Z_{eu} , , $g(I_0)$]

{5 to 100 mg m⁻³}

{0 to 2 div. d⁻¹}

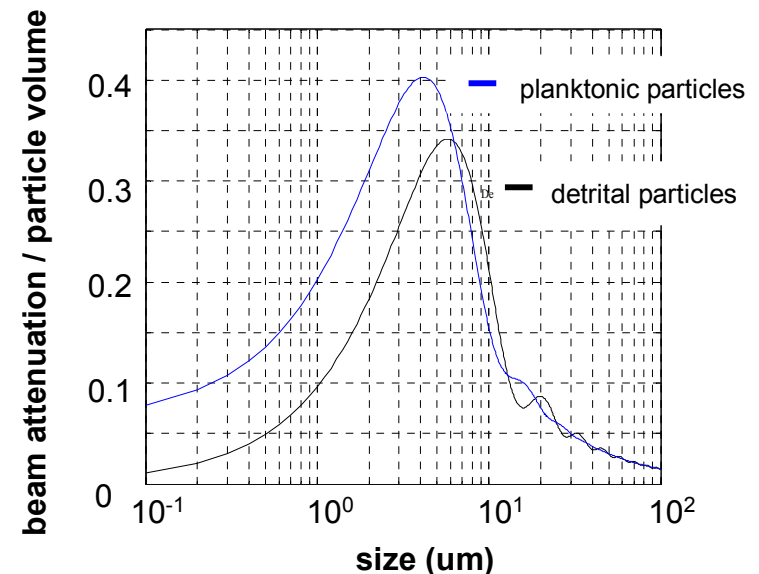
14 Steps

- (1) Light scattering 🛎 # particles
- (2) Common measures of scatter:
particulate beam attenuation (c_p)
particulate backscatter coeff. (b_{bp})
- (3) Open—ocean particle size spectra relatively conserved ➡

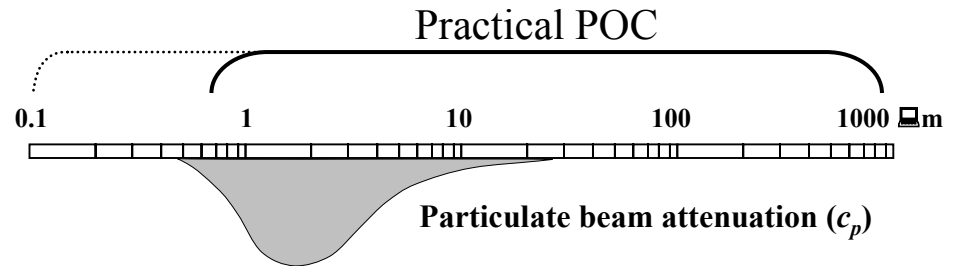


- (4) c_p is:

- dominated by particles in the phytoplankton size domain ➡
- sensitive to inorganic particles in the 0.5 to 20 μm size range
- insensitive to submicron bacteria, 'pure water' values are measured when 10^6 bacteria are present



(5) Particle # $\rightarrow c_p$ per particle



(6) c_p covaries with POC

Bishop, J.K.B. 1999. *Deep-Sea Res. I.* 46:353-369.

Bishop, J. K. B., S. E. Calvert and M. Y. S. Soon (1999) *Deep-Sea Res. II* 46:2699-2733.

Claustre, H., A. Morel, M. Babin, C. Cailliau, D. Marie, J-C Marty, D. Tailliez and
D. Vaultot (1999) *J. Geophys. Res.* 104:3401-3422.

Gardner, W. D., I. D. Walsh and M. J. Richardson (1993) *Deep-Sea Res. II.* 40:171-195.

Gardner, W. D., S. P. Chung, M. J. Richardson and I. D. Walsh (1995) *Deep-Sea Res. II.* 42:757-775.

Loisel, H. and A. Morel (1998) *Limnol. Oceanogr.* 43: 847-858.

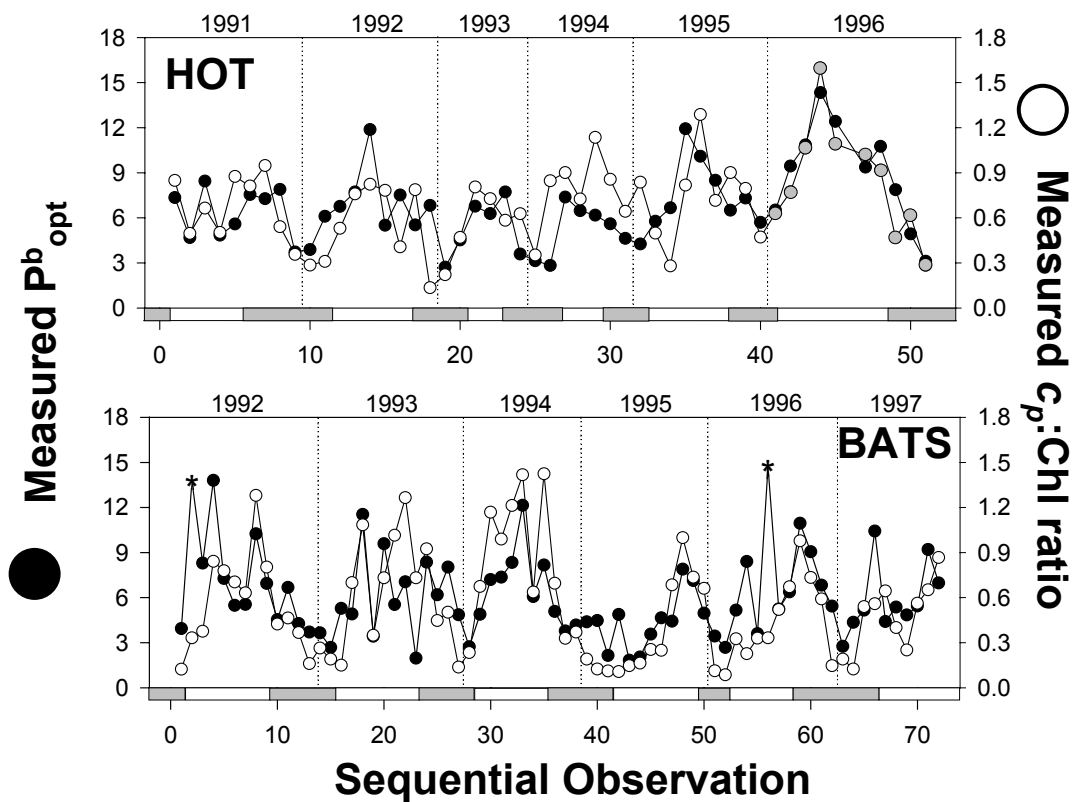
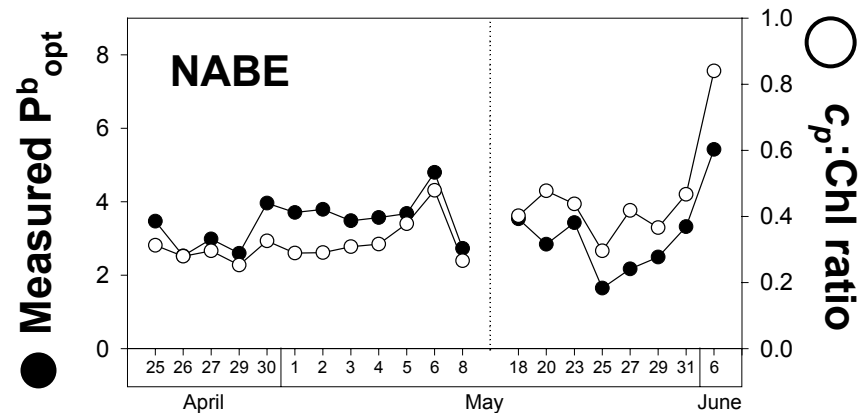
Walsh, I. D., S. P. Chung, M. J. Richardson and W. D. Gardner (1995) *Deep-Sea Res. II* 42:465-477.

(7) *Therefore* – c_p should be a better estimate of phytoplankton carbon than POC

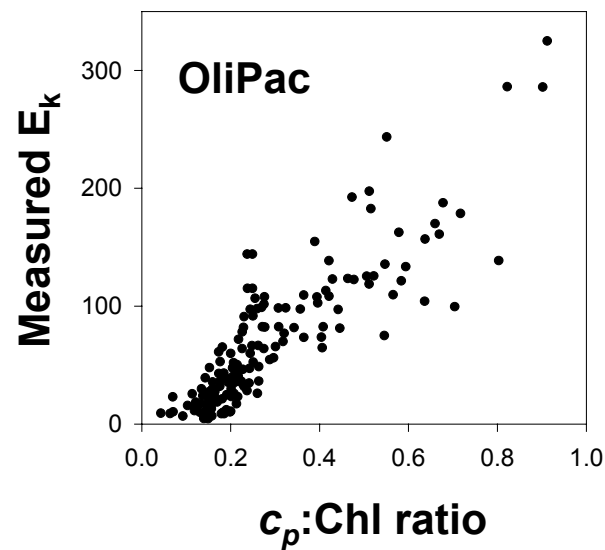
(8) *If true* – then the ratio of c_p :chlorophyll should track phytoplankton C:Chl, which is a common index of physiological variability that shares environmental dependencies with P_{opt}^b and E_k

(9) c_p :chl tracks changes in P^b_{opt} and E_k

Behrenfeld & Boss (2004) *Deep-Sea Res.*
50:1537-1549

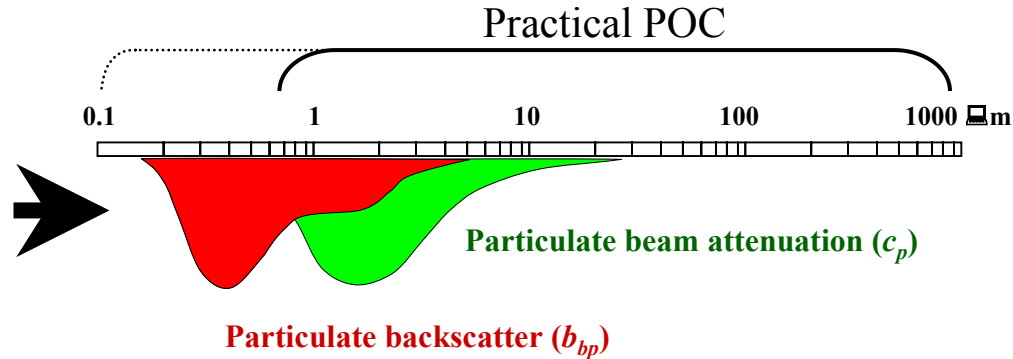


NABE Time Series



(10) Ocean color inversion algorithms give us b_{bp} , not c_p

(11) Mie calculations indicate that b_{bp} is dominated by submicron particles, but in field populations b_{bp} likely has a significant tail in the phytoplankton size domain.



(12) Satellite b_{bp} covaries with POC

Loisel, H., E. Bosc, D. Stramski, K. Oubelkheir, and P.-Y. Deschamps (2001)

Geophys. Res. Lett., 28:4203-4206.

Stramski, D., R. A. Reynolds, M. Kahru, and B. G. Mitchell (1999)

Science, 285:239-242.

(13) *Therefore* – b_{bp} should be a better estimate of phytoplankton carbon than POC

(14) *If true* – then the ratio of chlorophyll: b_{bp} should track phytoplankton Chl:C

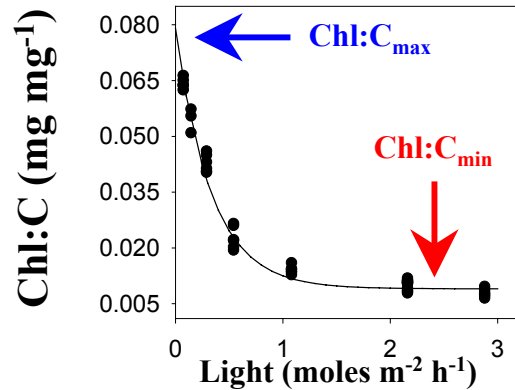
and thus phytoplankton growth rates (💻)

{once a correction of the background bacterial population is accounted for}

Chl:C physiology

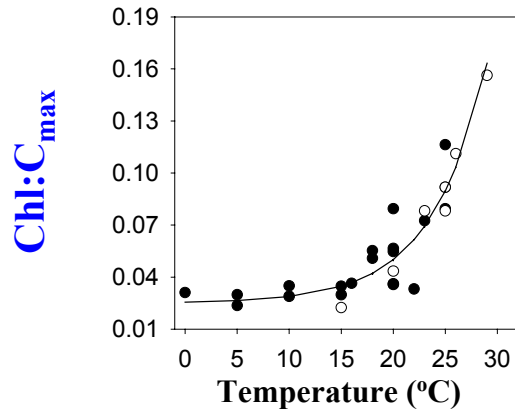
Laboratory

3 primary factors



← Light

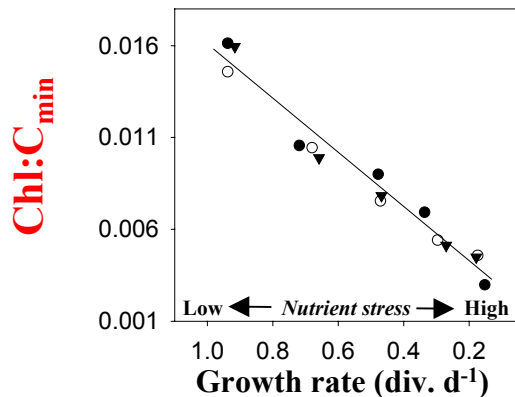
Dunaliella tertiolecta
20 °C
Replete nutrients
Exponential growth phase



← Temperature

Geider (1987) *New Phytol.* **106**: 1-34

16 species
● = Diatoms
○ = all other species

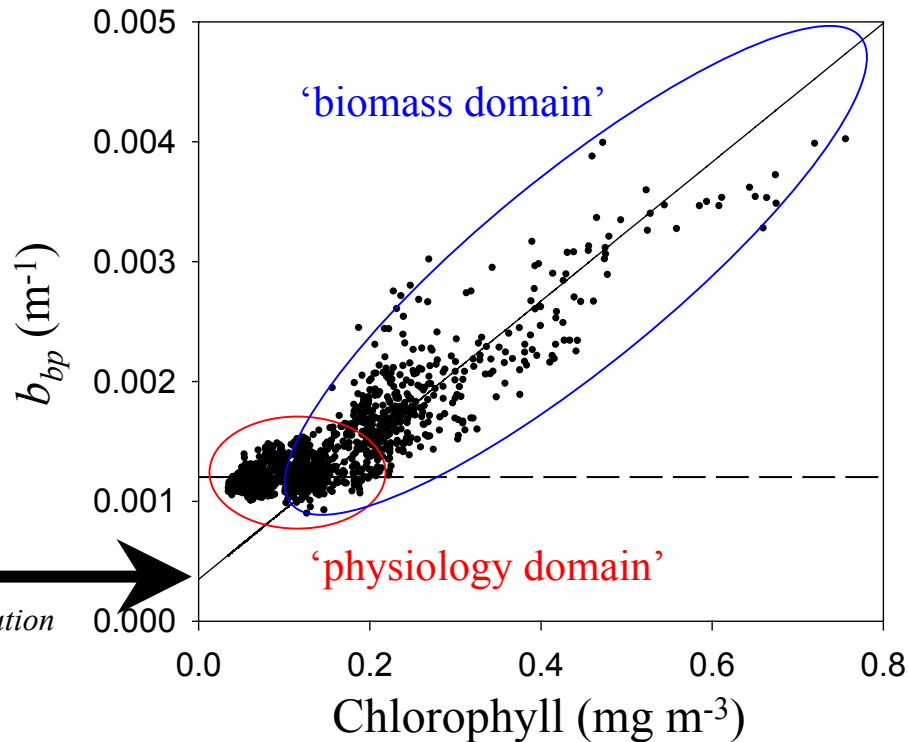
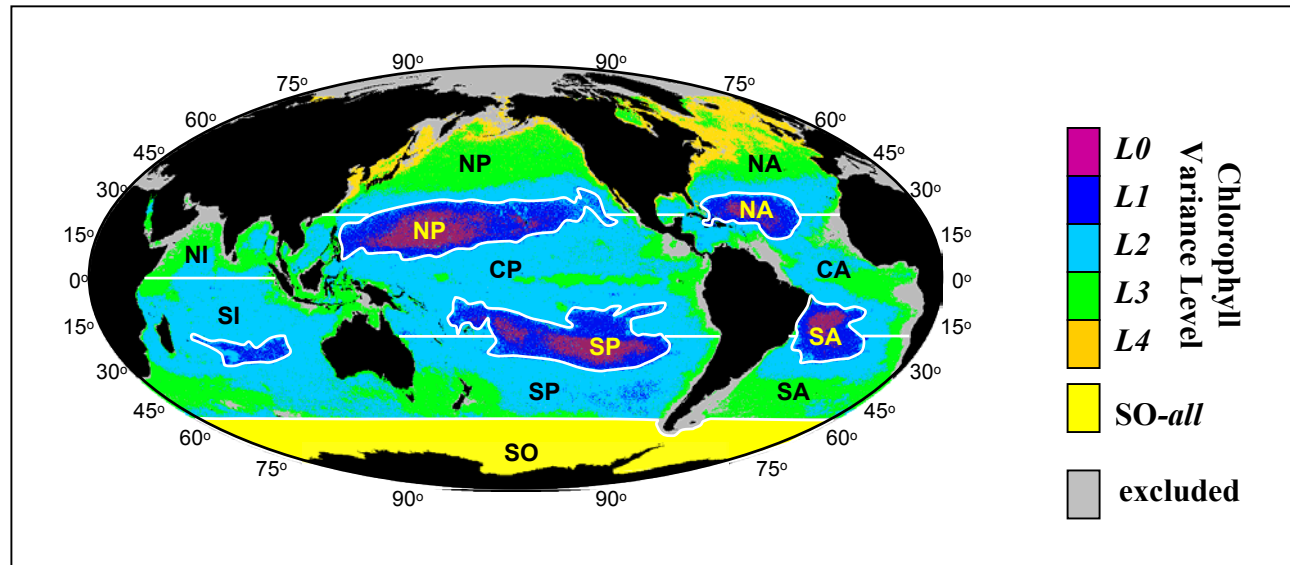


← Nutrients

Laws & Bannister (1980)
Limnol. Oceanogr. **25**: 457-473

Thalassiosira fluviatilis
● = NO₃ limited cultures
○ = NH₄ limited cultures
▼ = PO₄ limited cultures

28 Regional Bins



$$C = \text{scalar } \uparrow (b_{bp} - \text{intercept})$$

$$= 13,000 \uparrow (b_{bp} - 0.00035)$$

$$\text{Intercept} = 0.00017 \text{ m}^{-1}$$

Stramski & Kiefer (1991) *Prog. Oceanogr.* 28, 343-383

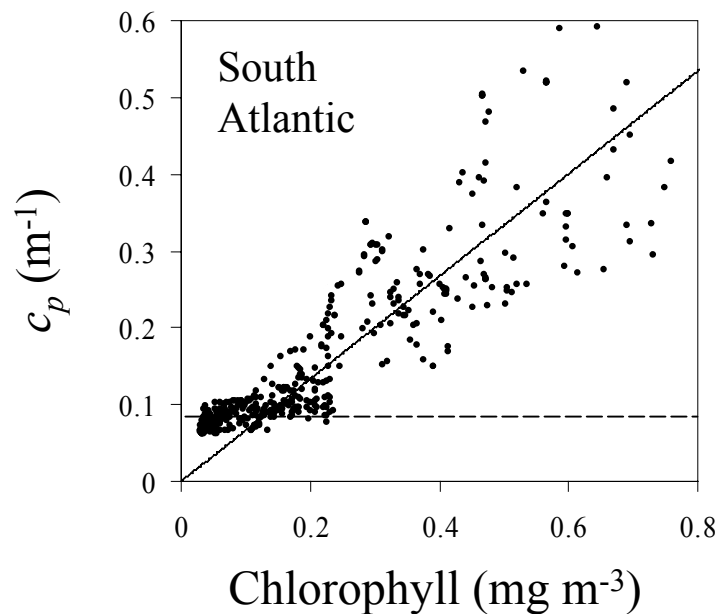
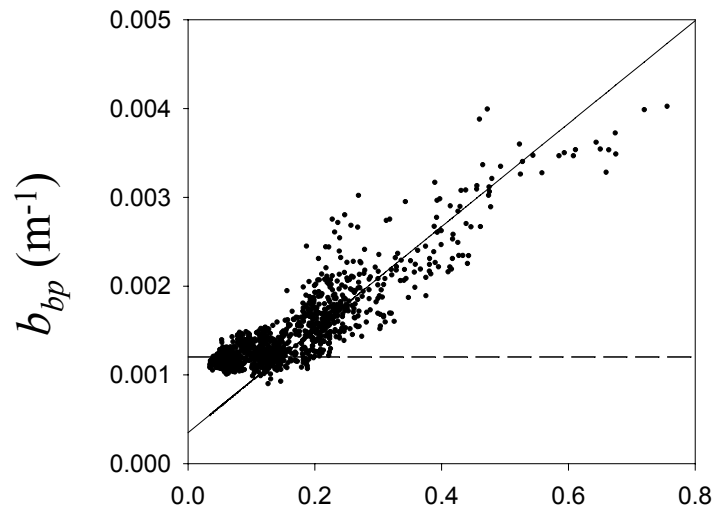
Cho & Azam (1990) *Mar. Ecol. Prog. Ser.*, 63, 253-259

$$\text{Phytoplankton Carbon} = 25 - 35\% \text{ POC}$$

Eppley et al. (1992) *J. Geophys. Res.*, 97, 655-661

DuRand et al. (2001) *Deep-Sea Res. II*, 48, 1983-2003

Gundersen et al. (2001) *Deep-Sea Res. II* 48, 1697-1718



Satellite b_{bp} vs field c_p

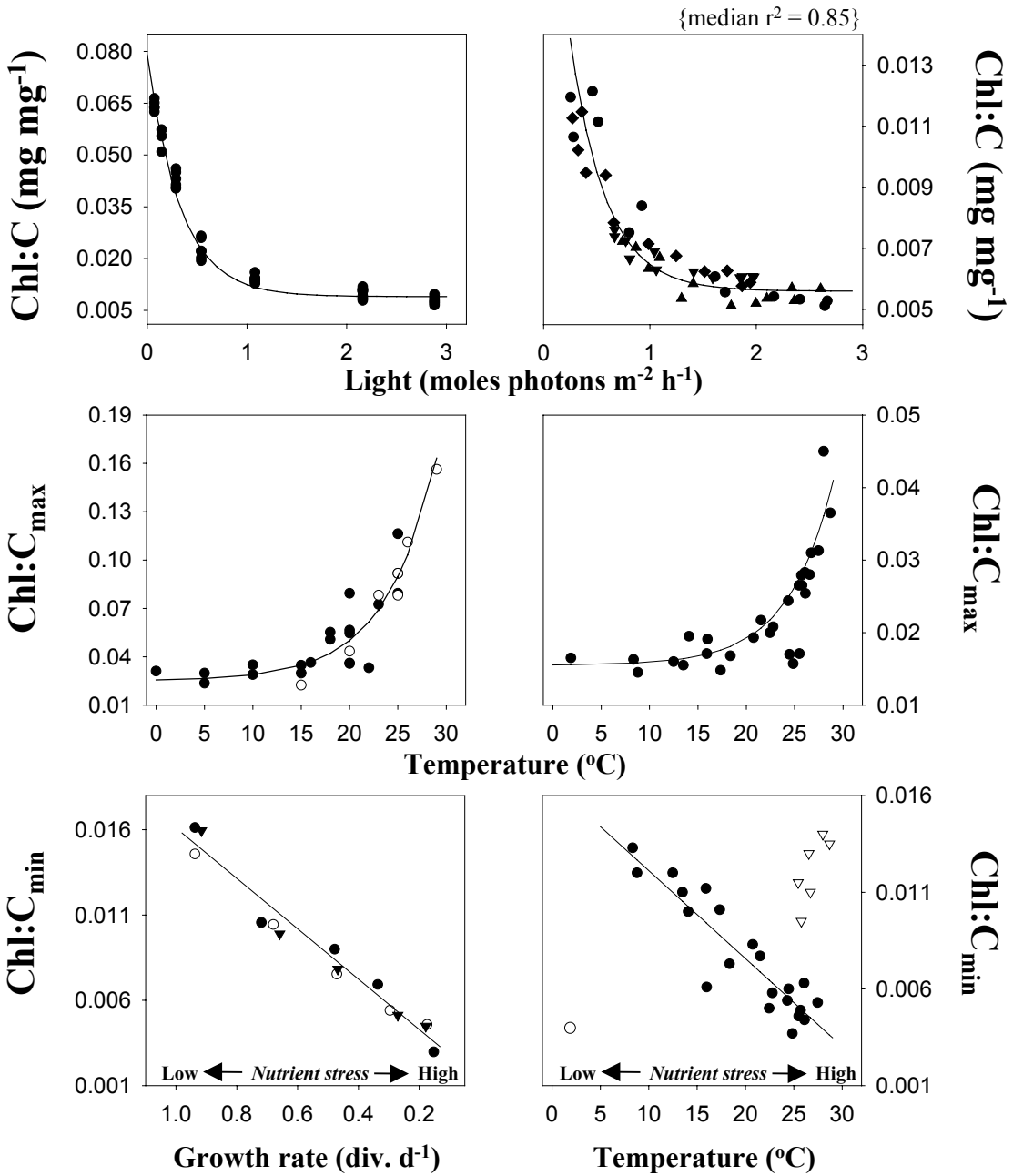
- Same bilinear pattern
- Anticipated intercept

$$b_{bp} > 0$$

$$c_p = 0$$
- Ratio of slopes = 0.0008

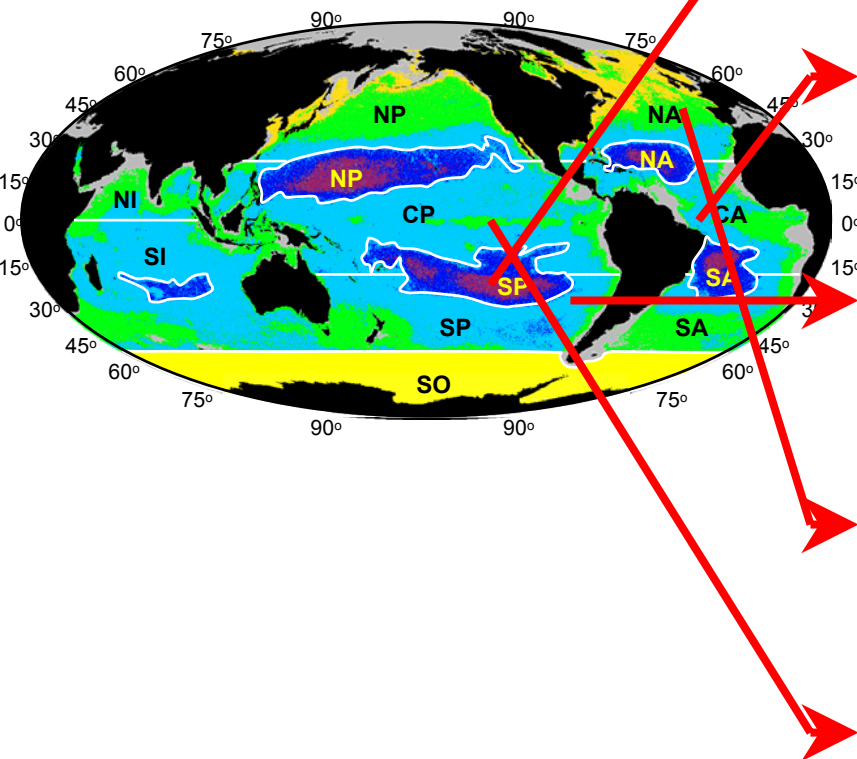
$$\{\text{phytoplankton} = 0.0005 - 0.0009\}$$

Laboratory

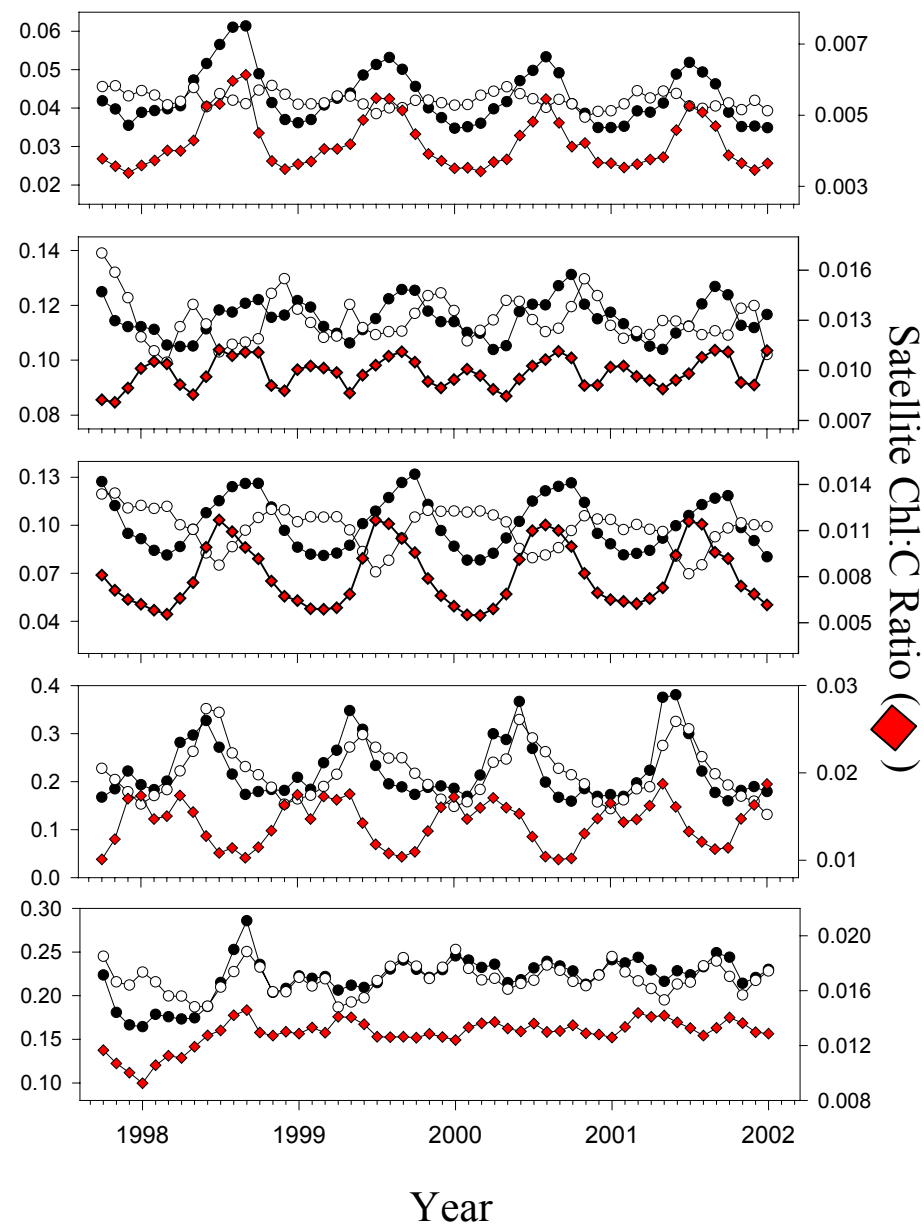


Space

The 5 basic patterns



Phytoplankton Chlorophyll (●) and scaled Carbon (○)



$$\text{Growth rate (🖥️)} = \text{🖥️}_{\text{max}} \quad \text{👉} \quad f(\text{nuts, temp}) \quad \text{👉} \quad g(\text{light})$$



2 divisions d⁻¹

Banse (1991)



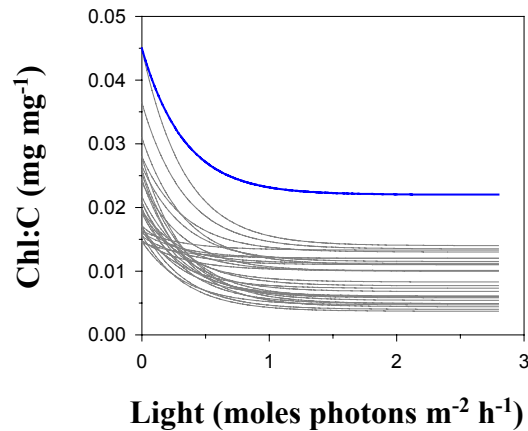
range = 0 to 1



range = 0 to 1



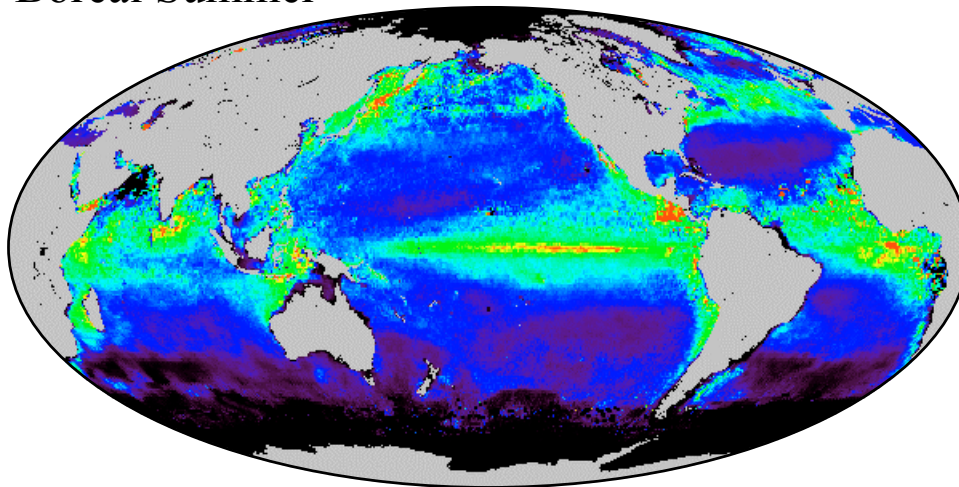
$$1 - \exp^{-3\text{light}}$$



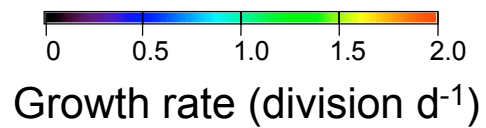
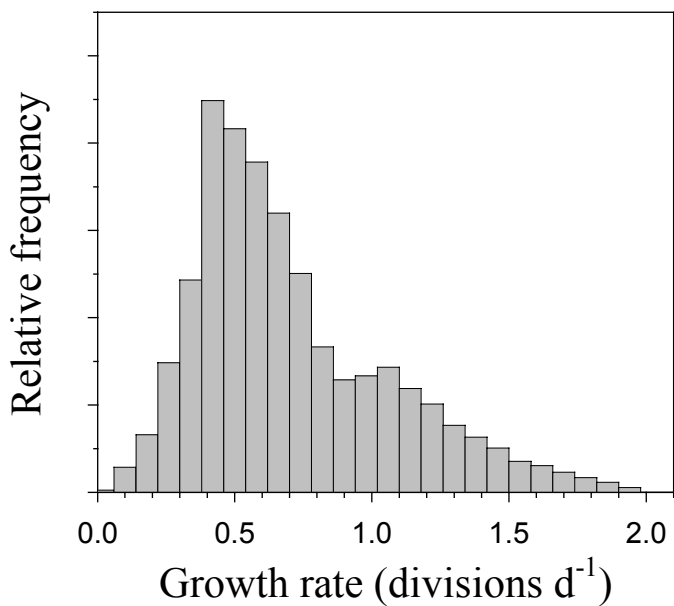
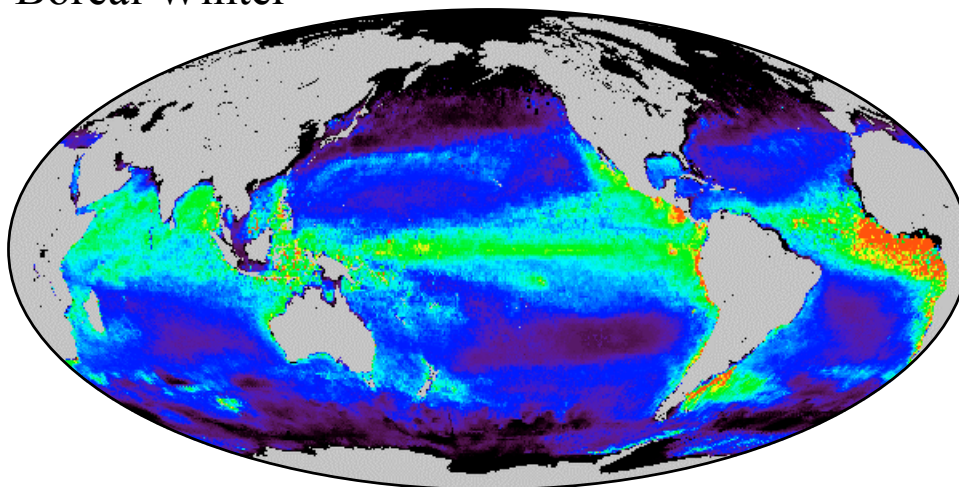
Chl:C_{sat}

(Chl:C_{N,T-max}) Light

Boreal Summer



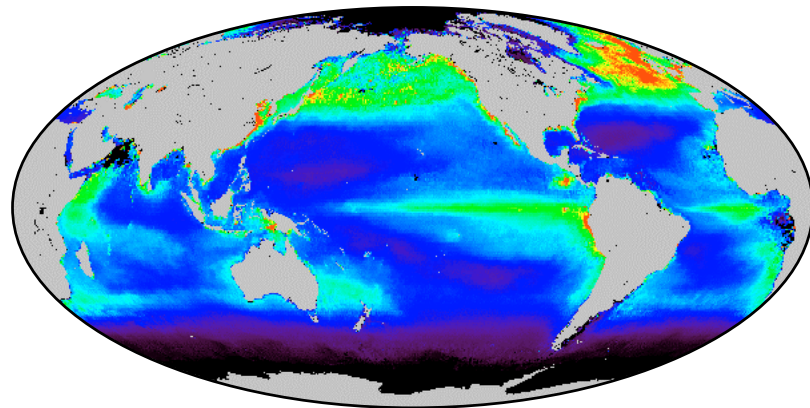
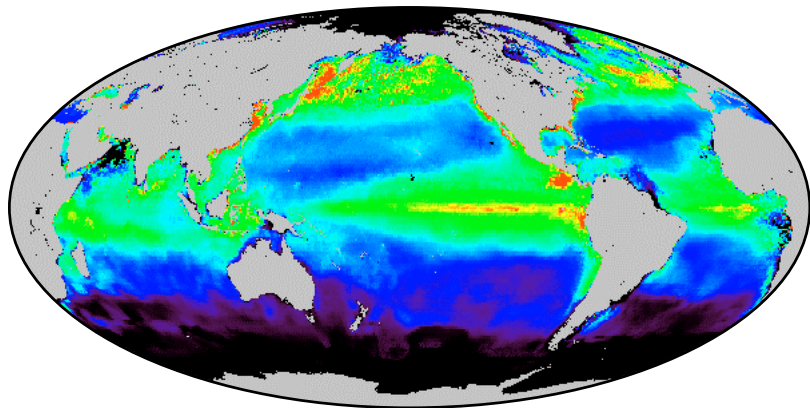
Boreal Winter



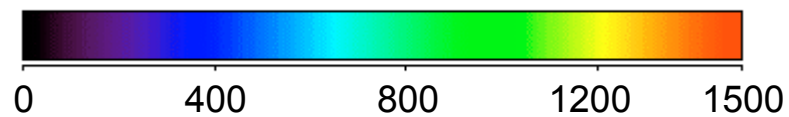
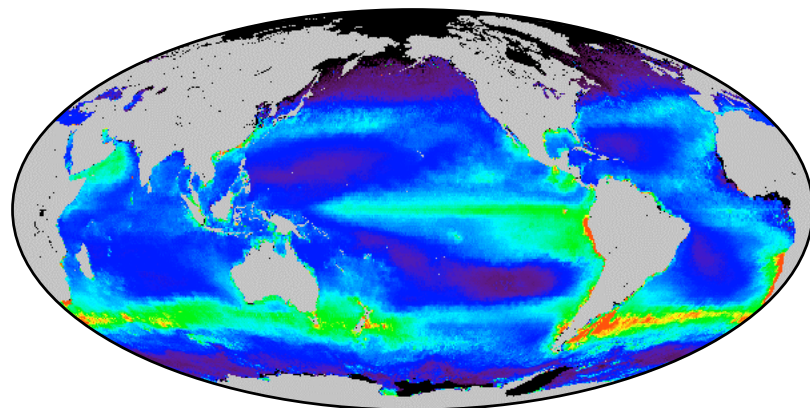
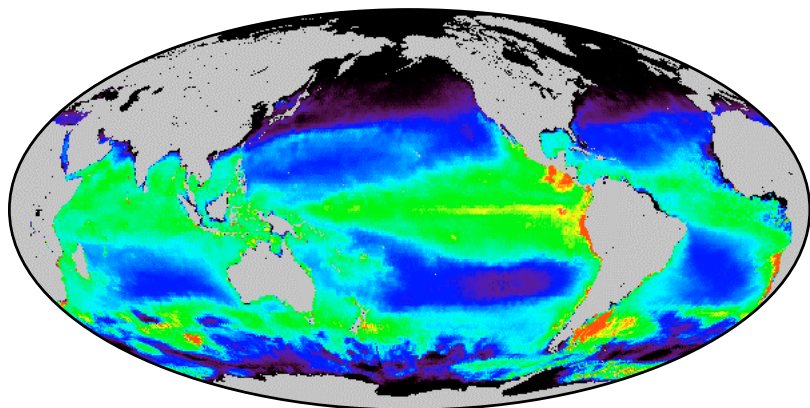
Carbon based

Chlorophyll based

Boreal Summer



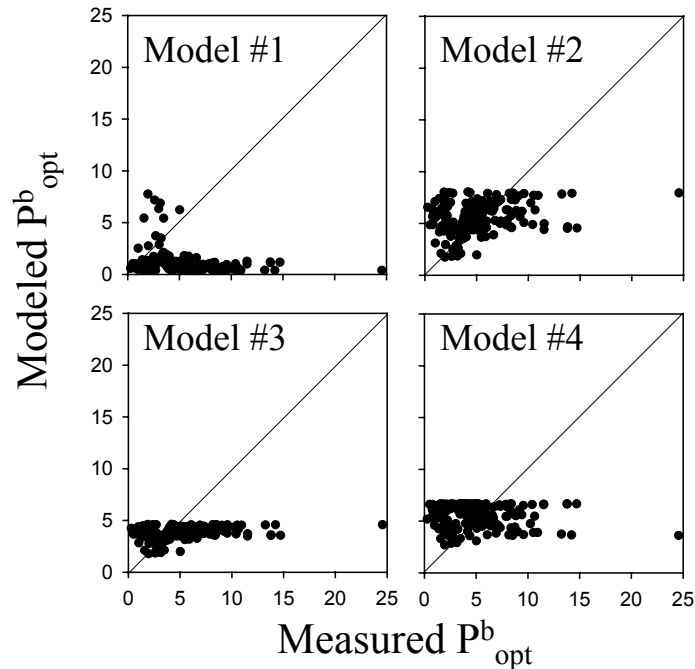
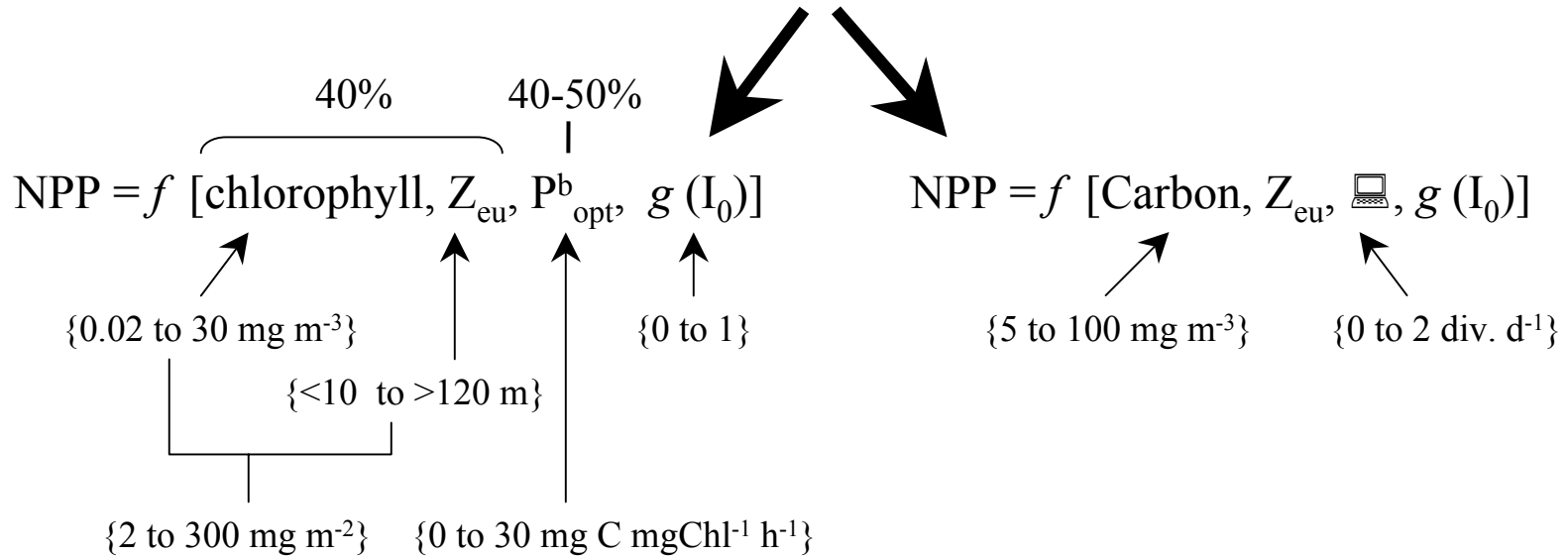
Boreal Winter

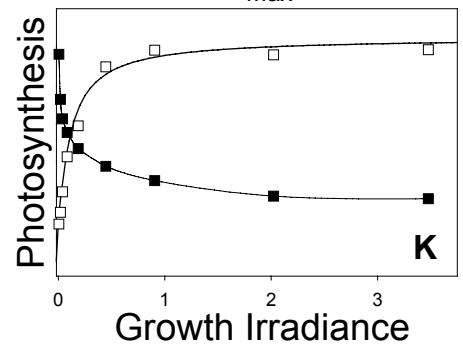
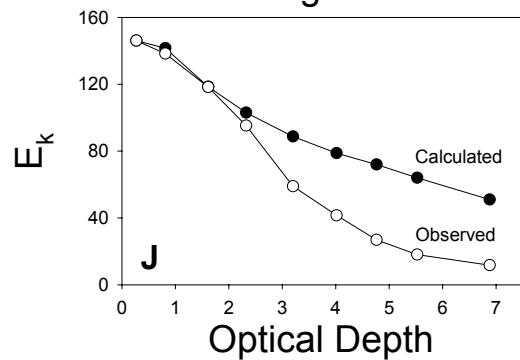
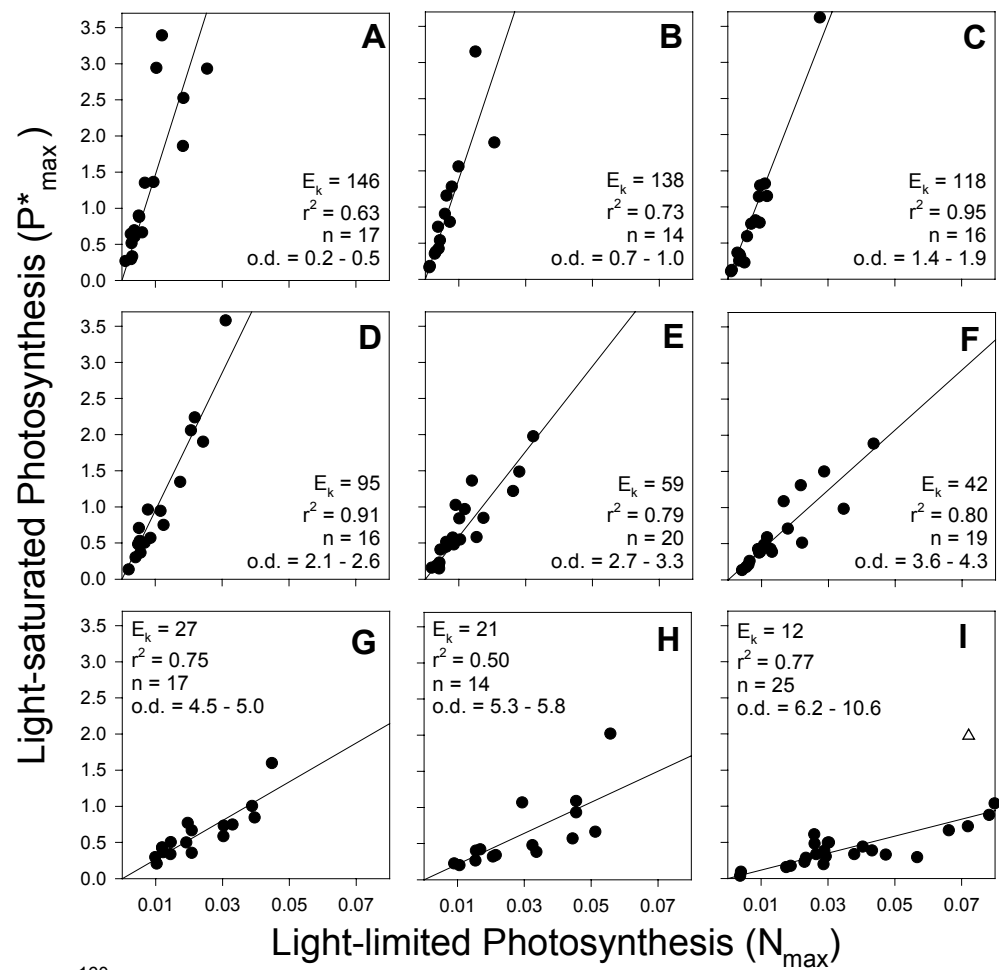


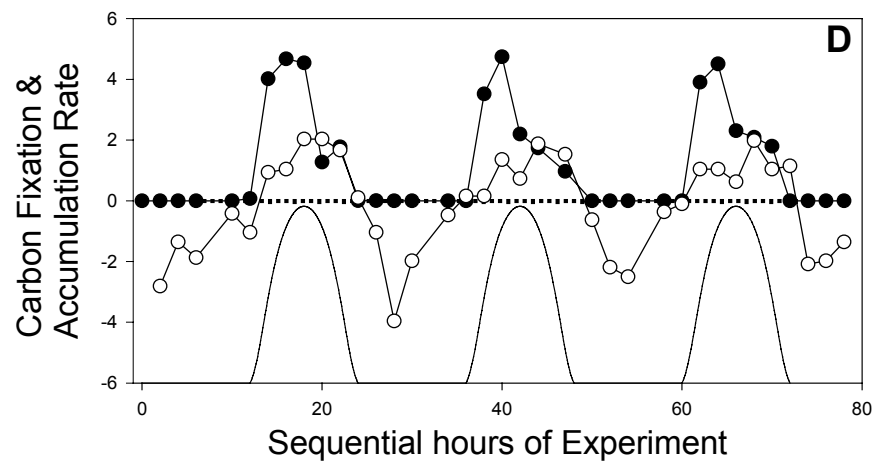
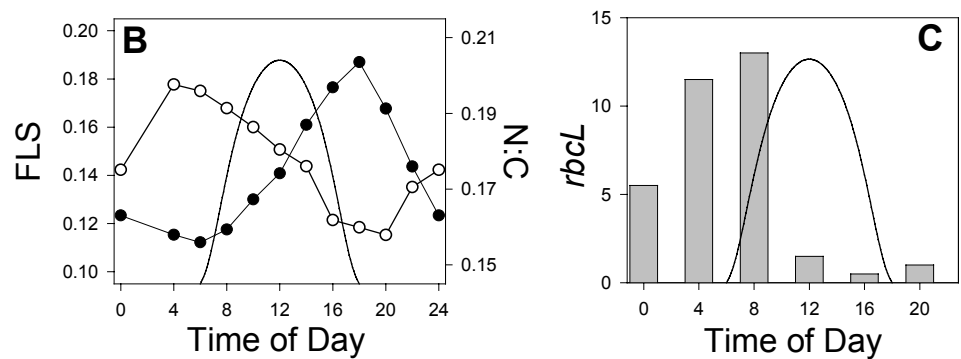
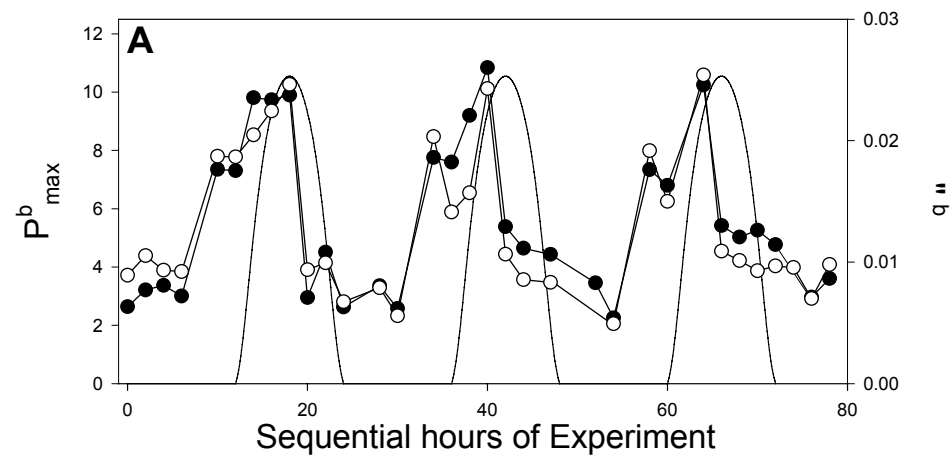
Net Primary Production (mg m⁻²)

Backup Slides

Net primary production (NPP) = f [biomass, physiology, $g(I_0)$]







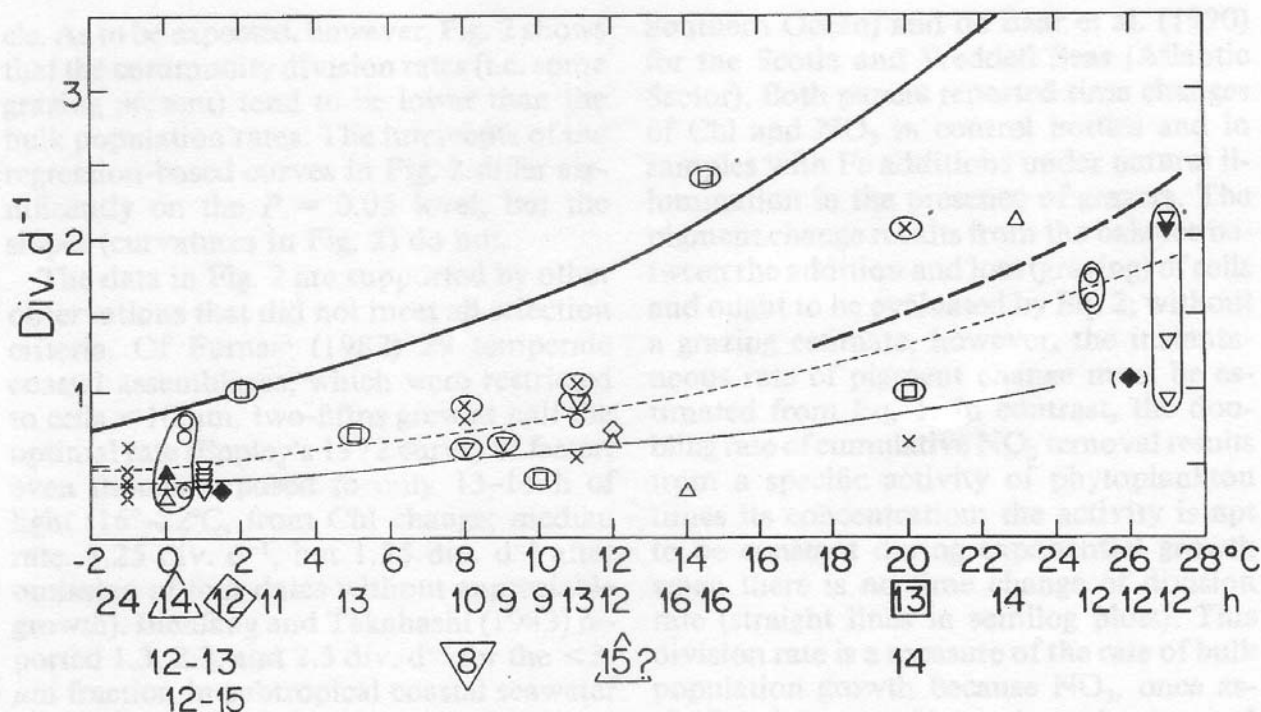


Fig. 2. Division rates for natural marine plankton in enclosures as dependent on temperature, with the "standard curve" for optimal growth (Eppley 1972) and half of this maximum (broken curve) added. Estimated daylength (h) shown below temperature scale. Filled symbols—two stations. Dashed line (from regression, see text) and circled or enclosed symbols—bulk population division rates [results without grazing, from dilution method, or Chl labeling (the latter for Welschmeyer and Lorenzen 1984; Downs and Lorenzen 1985; Laws et al. 1984)]; thin line and other symbols—community division rates (from Chl increases). From left: $\times$ —Spies 1987; Δ , $\triangle$ —Paranjape 1987 (samples with excess NO_3); $\circ$ —Taylor and Haberstroh 1988; ∇ —Sakshaug and Holm-Hansen 1986 (highest values at intermediate irradiances); $\blacklozenge$ —Kuiper et al. 1983 (examples of high rates, No. 24, 25; temperature from Jahnke et al. 1983); $\square$ —Gifford 1988; $\times$ —Welschmeyer and Lorenzen 1984 [table 3]; $\otimes$ —highest values except for March (second highest); $\times$ —means of two highest values for each date; temperatures from original log books]; ∇ —Downs and Lorenzen 1985 [table 3; means of two highest values for November, January and April, when NO_3 and incident irradiance were high (from original log books)]; $\circ$ —Wilkerson and Dugdale 1987 (No. 58B 2); $\diamond$ —Riemann et al. 1988 (enclosure C, 12–19 September); $\triangle$ —McAllister et al. 1961 (days 4–14), Antia et al. 1963 (days 8–15), both at low irradiance; $\triangle$ —Eppley et al. 1971 (from figure 2, 11–13 July average of NO_3 and NH_4 additions; "20°–25°C"); $\circ$ —Landry et al. 1984 (temperature estimated); $\blacklozenge$ —Cullen et al. 1992 [mean of several stations, under low irradiance (moved from Fig. 3, not used in regression)]; ∇ , ∇ —Laws et al. 1984 (without offshore station F).